#### Geochemistry and Stable Isotopes of the Rift-Related Quartz-Adularia-Type Gold-Silver Mineralization, Bergama, Izmir, Turkey

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80 g/t Au and 91 g/t Ag

# Presentation will be done under the following headings:

- Location/purpose of study
- Geological setting
- Vein system/mineralogy
- Hydrothermal alteration
- Hydrothermal Geochemistry
- Fluid inclusion
- Stable isotope studies
- Discussion

#### Location

# Ovacık-Narlica epithermal gold deposits are located in the west Anatolian extensional province.



#### Oligo-Miocene subaerial volcanics hosting several epithermal Au deposits

Kuçukdere LS, 60 km N, 1.4 mt @6.5 g/t Au.
Efemcukuru LS, 100 km S, 3.1 mt @ 14.6 g/t Au.
Kisladag HS, 200 km E, 276 mt @1.2 g/t Au.



# Geological Setting

#### **Regional Geology**

- A major granodiorite intrusive, surrounded by calc-alkaline volcanic rocks such as andesite, dacite and rhyodacite.
- Tectonics: A strong compression from Eocene to Miocene, causing the generation of anatectic granitic melts.



### Local Geology

 The Ovacik-Narlica lithologies consist mainly of K-rich subaerial andesite and dacite with minor latite andesite

 Ovacik and Narlica veins are located next to graben margin.



# Characteristics of vein system

Ovacik vein system consists of four epithermal quartz veins; of these M and S veins are economically significant, both having a total mapped strike length of 900 m.



#### Vein and Breccia Textures

Main textures are: Colloform/Crustform banding, occurring in veins and breccia clasts (3 m @ 28 g/t Au)



#### Coarse, bladed carbonate replacement textures occurring as distinct bands or infilling vugs (0.08 g/t Au)



Matrix-supported fluidized (milled) breccia consisting of quartzadularia vein material with common crustiform banding, bladed calcite (2 m @ 39.2 g/t Au)



#### Clast-supported crackle (shatter) breccia

 consisting of qurtzadularia vein material with common crustiform banding (1 m @ 24.5 g/t Au)



#### Vein Mineralogy Mineral paragenesis

# An approximate sequence of mineralization is:

- Quartz + Pyrite + Illite + Smectite >
- Quartz + Pyrite + Illite + Smectite + Calcite + Adularia + Arsenopyrite + Chalcopyrite + Electrum >
- Chalcopyrite + Bornite+ Fahlerz + Galena + Electrum + Argentite +Acanthite + Marcasite + Pyrite >
- Kaolinite + Marcasite + Native Au + Native Ag + Covellite

Mineral	Pre-ore stage	Ore stage I	Ore stage II	Post ore/Supergene
Quartz		-		
Adularia		-		
Calcite*	-		-	
Illite/Sericite				
Illite/Smectite			_	
Smectite			-	
Kaolinite				
Rutile	-		-	
Pyrite				
Marcasite			-	
Arsenopyrite				
Chalcopyrite				
Bornite				
Fahlerz group				
Sphalerite				
Galena				
Electrum		-		
Native gold				
Argentite				
Acanthite				
Native silver				
Covellite				
Fe oxides				
*Mainly replace	ed by quartz			

## **Hydrothermal Alteration**

- Silica alteration enveloped by illite/sericite, minor mixed layer illite/smectite.
- Adularia occurring as vein material and as mass replacement of andesite.
- Ovacik deposit forming at 200 °C evidenced by the presence of illite, adularia, illite/smectite whereas Narlica deposit forming at higher temperatures > 200 °C evidenced by the occurrence of highly crystalline quartz.



# Fluid Inclusion

- Homogenization temperatures of quartz veins range from 147-298
   °C, with an average mode of 190 °C.
  - The presence of vapor-rich inclusions, variable liquidto-vapor ratios and large ranges in temperatures (indicate phase separation /boiling).
- Ice melting temperatures (*T*m) range from -0.4 to 1.2 degree °C, indicating a very dilute salinity less than 2 eq. Wt.%.





#### Hydrothermal Geochemistry Wallrock Geochemistry

K, Rb and Cs show twofold enrichments caused by formation of adularia or illite/sericite.

- Ba, Th, U, Nb, Sr, Nd, Zr, Ti, Y, Cr, Ni, Cu and Pb depletions are generally <three-folds whereas Sr is depleted by a factor of 5.
- REE display a flat pattern.



#### Hydrothermal Geochemistry Quartz Veins

- Ba, Th, U, Nb, La, Ce, Sr, Nd, Zr, Ti, Tb, Y and Cr depleted by factors of 4 to 23 with negative anomalies for Th at 13 and for Sr at 23.
- Mg, Ca and Na are reduced by factors of 48, 60 and 84, respectively.
- LREE (La-Eu) values are reduced by factors of 7 to 15.
- HREE showing a linear trend from Gd to Yb appears with depletion factors of 4 to 9.



#### Hydrothermal Geochemistry Chondrite-Normalized REE

Evidence for remobilization of REE at Ovacik-Narlica is provided by their consistent decrease from fresh volcanics through clay and adularia-altered wallrock to quartzadularia veins.



#### Hydrothermal Geochemistry Au and Other Metals Associations

 Moderate to strong correlations occur between Au and Ag, Pb, Zn, Cd, Cu, Sb indicating similar mineralizing event.

Very weak correlation between Au and As may suggest possible introduction of As and Au in different phases.



### Hydrothermal Geochemistry Distribution of Au,

- Very high gold values are confined to M vein and decease remarkably eastward (S vein).
- High silver values are confined to upper portion of M vein.
- Antimony values are confined to upermost portion of the two veins



## Stable Isotope Oxygen Isotope

δ<sup>18</sup>O-quartz values in M and S veins from Ovacik deposit range from +9.46 to +15.71 %, and calculated δ<sup>18</sup>O-H<sub>2</sub>O values range from -2.86 to +3.51 ‰ (variation.:6.4)

 Average δ<sup>18</sup>O-H<sub>2</sub>O value (-0.64 ‰) of M vein is relatively lower than that of S vein (+1.1‰).

 Calculated δ<sup>18</sup>O-H<sub>2</sub>O values at Narlica range from -3.90 to -6.27 ‰ (variation.:2.4).



## Stable Isotope Oxygen Isotope

- Oxygen isotope distribution at deeper levels are relatively homogenous and they exhibit significant variations at higher levels.
- Fluid inclusion water at Ovacik have isotopic compositions (Ave+0.6‰) which are enriched in heavy δ<sup>18</sup>O as compared to present-day meteoric (Ave-5.4‰) and geothermal water (Ave-6.8 ‰).



EASTING (m)



#### Stable Isotope Hydrogen Isotope

Ovacik-Narlica quartz samples yield  $\delta D_{H_2O}$ isotopic compositions of -92 to -117 ‰ with an average of 102 ‰.

 The quartz has low δD-H<sub>2</sub>0 isotopic values as compared to recent active geothermal waters (Ave. δD: -42 ‰).



## Stable Isotope Sulfur Isotope

 δ<sup>34</sup>S-H<sub>2</sub>O compositions of S and M veins range between -2.98 and +2.5 ‰ (Ave.:+0.44 ‰) and -2.98 and +4.4 ‰ (Ave.:+1.1).

 $δ^{34}S_{H_20}$  values of sulfides are clustered near zero (ave. for M and S vein:+0.8 ‰), indicating presence of H<sub>2</sub>S as probable dominant species.



### Stable Isotope

![](_page_23_Figure_1.jpeg)

- Fluids of meteoric origin appear to be dominant with lesser magmatic contribution at Ovacik deposit.
- LS deposits elsewhere in the world commonly show an O-shift from local meteoric water values, whereas high-grade ore samples as in the Comestock Lode epithermal gold deposit have both an O- and D-isotopic shift from local meteoric water, likely caused by magmatic water addition (Hedenquist and Lowenstern,1994)
- Such significant shift at Ovacik deposit can not only be caused by boiling of meteoric water.

![](_page_24_Figure_0.jpeg)

- Kodera et al. (2005) suggested that boiling of 90 % of liquid would result in a maximum enrichment of 3.9 ‰ δ<sup>18</sup>O and a maximum relative depletion of 13 ‰ in δD.
- The variation in the samples reported at Ovacik (up to 6.4 ‰ for δ<sup>18</sup>O fluid and 25 ‰ for δD fluid) therefore, demonstrate that the end stages of open-system boiling and fractionation could not have been reached everywhere, but locally.

#### Stable Isotope/Mineralogy/Au Geochemistry

![](_page_25_Figure_1.jpeg)

- In M vein (Ovacik W), spatial association of highest Cu+Pb+Zn and Au values with depleted δ<sup>18</sup>O-H<sub>2</sub>O is noticed.
- The low δ<sup>18</sup>O-н<sub>2</sub>o (bulls-eye) may represents the locus of the highest water/rock ratio.
- Therefore, this may be considered as an important guide for further exploration in the region.

### Conclusions

- Gold deposits are of low sulfidation.
- Ore minerals are mainly electrum, acanthite, tetrahedrite (Cu<sub>12</sub>Sb<sub>4</sub>S<sub>13</sub>), rare pyrargyrite (Ag<sub>3</sub>SbS<sub>3</sub>), mostly in M vein.
- Temperatures of the formation of the alteration and mineralization are~ 200 °C at Ovacik and 260 °C, at Narlica, deduced from fluid inclusion and alteration minerals. Salinity is< 2 eq. wt. % NaCl.
- Enrichment of K, Rb, Cs (in Adularia, sericite) by a factor of 2 and depletion of Mg, Na, Ca, Sr by a factor of >5 in altered wall rock show that significant water/rock interaction occurred.
- Depletions of Sr by a factor of 23 and, of Mg, Ca and Na and by factors of 48, 60 and 84, respectively, in *mineralized quartz veins* points to a very significant water/rock interaction.
- LREE (La-Eu) concentrations in quartz veins were reduced by factors of 7 to 15, indicating mobilization of REE under conditions of significant fluid flow. This may be important guide/hint in exploring for potential Au-Ag fields.

# Conclusions

- Moderate to strong positive correlations (R=0.51-0.72) between Au and Ag, Pb, Zn, Cu, Sb suggest that they may be related to the same mineralizing event.
- Very weak correlations between Au-As and Ag-As indicate different mineralizing events, introduction of As and Au-Ag in different phases.
- Ore-forming hydrothermal fluids have δ<sup>18</sup>O-H<sub>2</sub>O values (Ave: -0.6‰), where <sup>18</sup>O is enriched compared with present-day geothermal water (-6.8 ‰, from hot springs and production wells).
- δ<sup>18</sup>O values in the Ovacik deposit have been shifted from meteoric (-5.4‰) to hydrothermal (-0.6‰), to more <sup>18</sup>O-rich compositions by water/rock interactions or probably by some magmatic water contributions.

## Conclusions

 In M vein, coincidence of high Cu+Pb+Zn and Au values with depleted δ<sup>18</sup>O-H<sub>2</sub>O isotope values (bulls-eye) may represent the locus of the highest water/rock ratio.

 A larger area with clay-adularia alteration (~30 km<sup>2</sup>) in the study area may be sampled for Oxygen Isotope analysis to identify the δ<sup>18</sup>O-depleted locations, which may further lead us to identify the potential fields for gold mineralizations.

![](_page_29_Picture_0.jpeg)